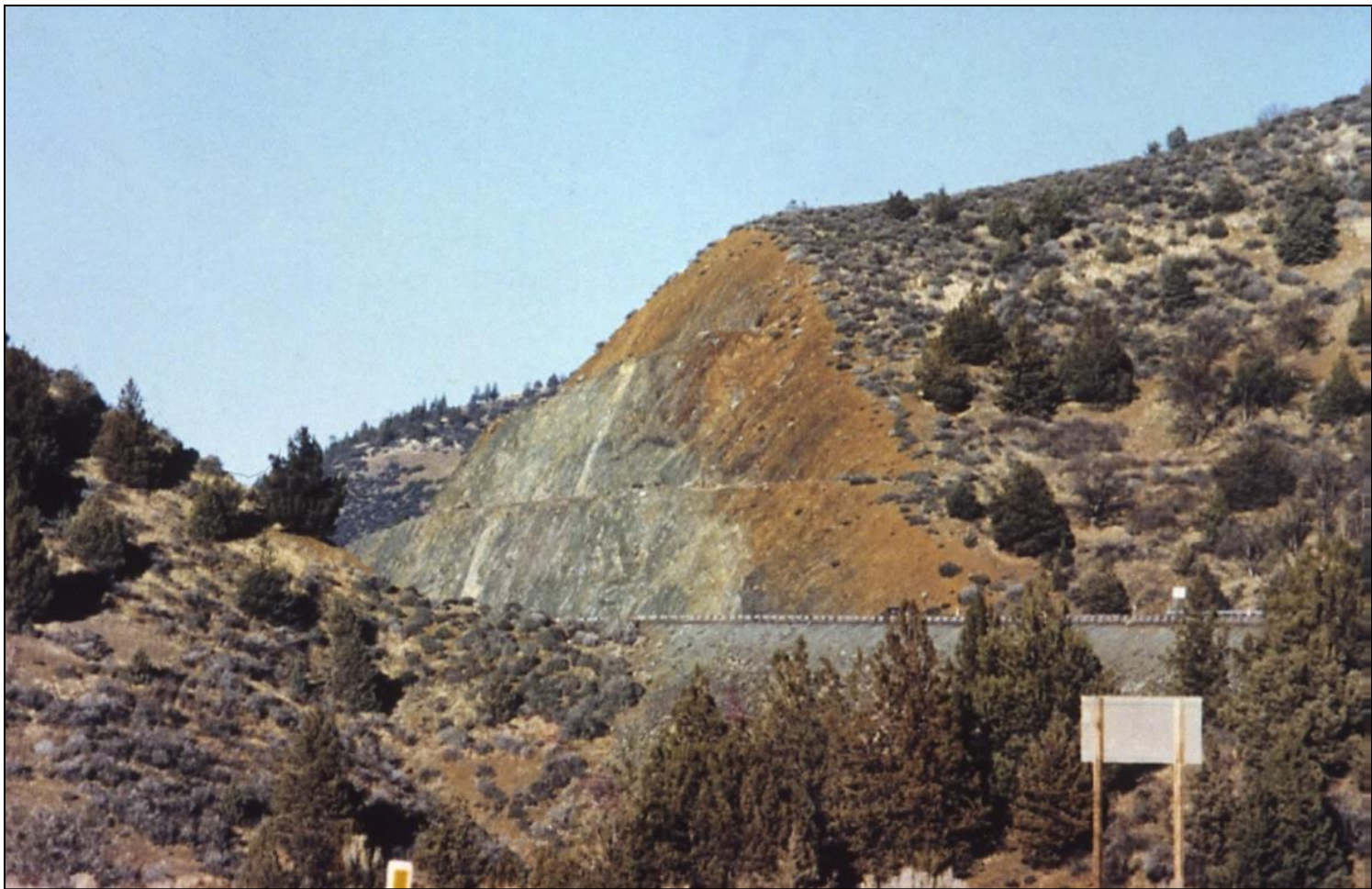


Part 3

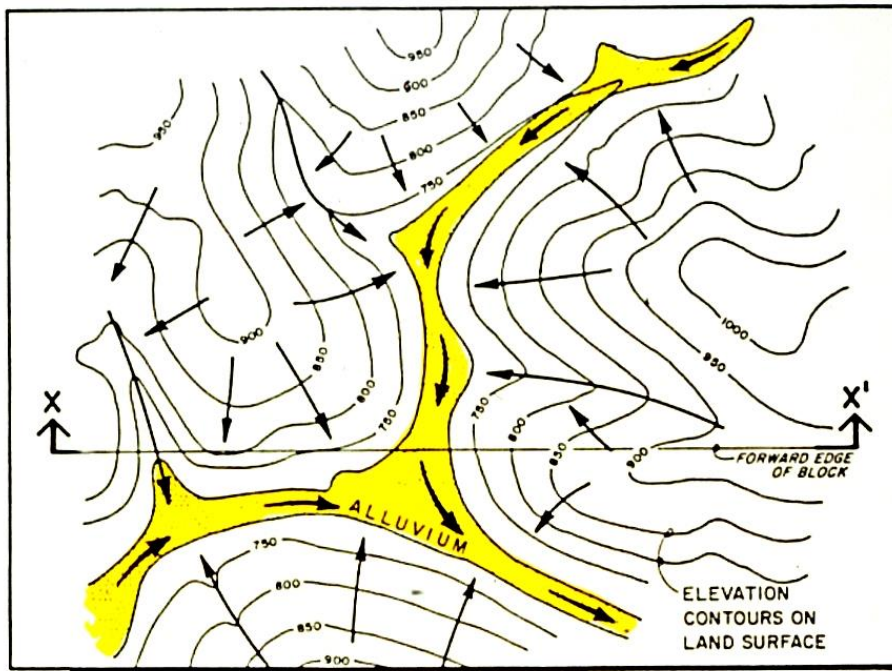
SHALLOW SEEPAGE IN THE WEATHERED BEDROCK CREEP ZONE



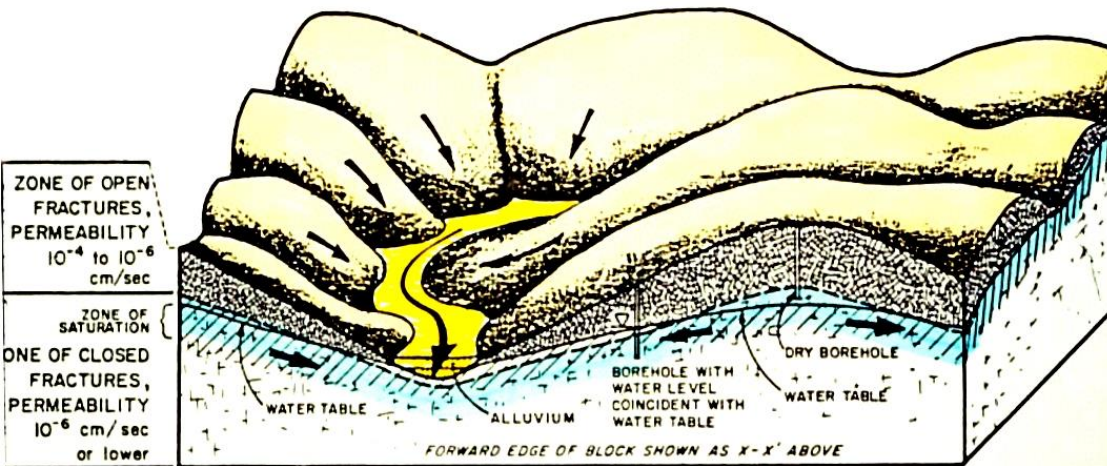
- Shallow **subsurface seepage** tends to percolate most easily through the weathered rock horizon, shown here as the reddish colored layer above the parent greenstone bedrock lying beneath.



- The bedrock creep zone lying beneath slopes is usually dilated and pervious, allowing ephemeral **seepage** to percolate through it.



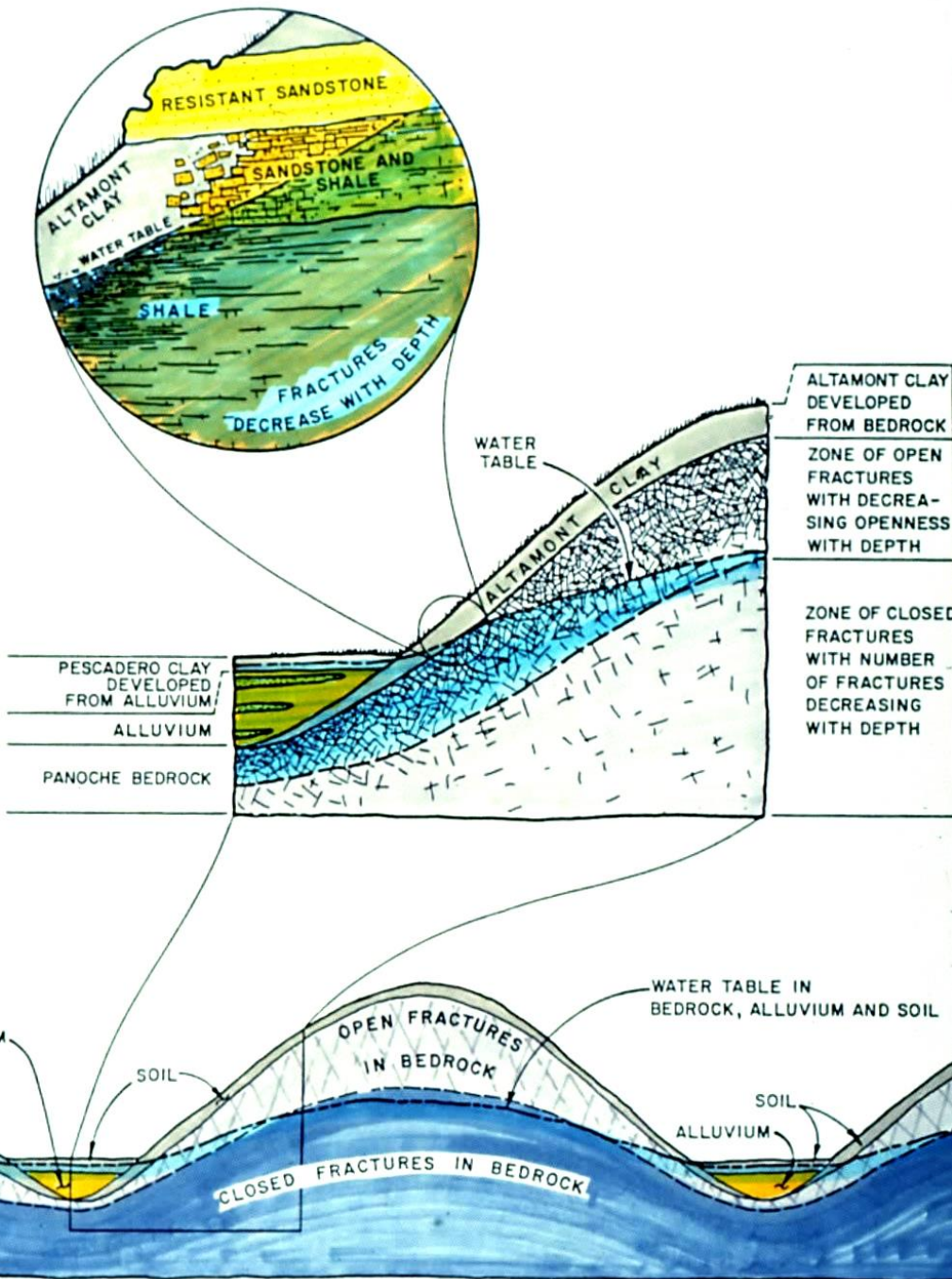
PLAN VIEW

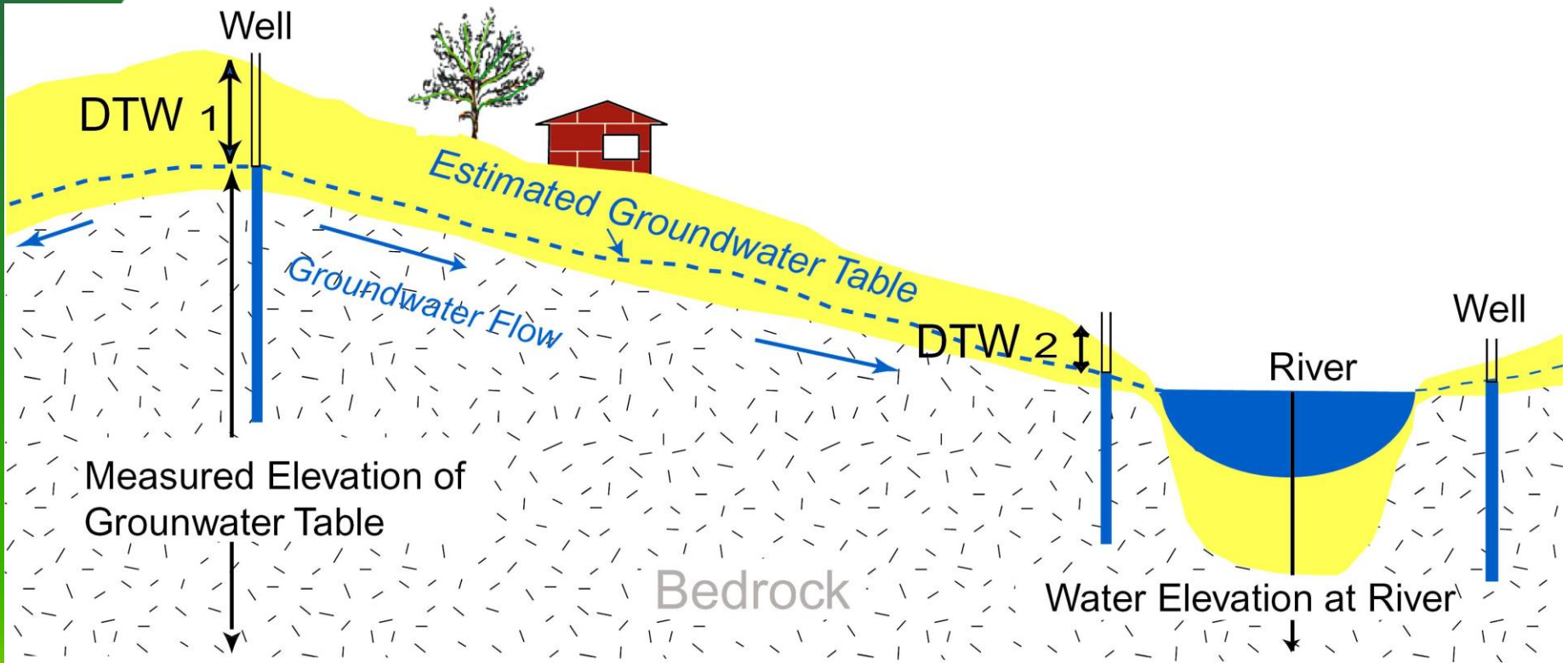


OBLIQUE VIEW

- **Seepage** tends to flow downslope, mimicking the surface topography
- The **groundwater table** usually lies within or close to the weatherized/oxidized zone, with flow towards the valley bottom, as shown here

- The bedrock creep zone is normally coincident with the partially saturated zone lying above the permanent **groundwater table**
- Seasonal fluctuations in groundwater levels and flowage usually occur within this zone, shown by the light blue areas above the dark blue at lower left

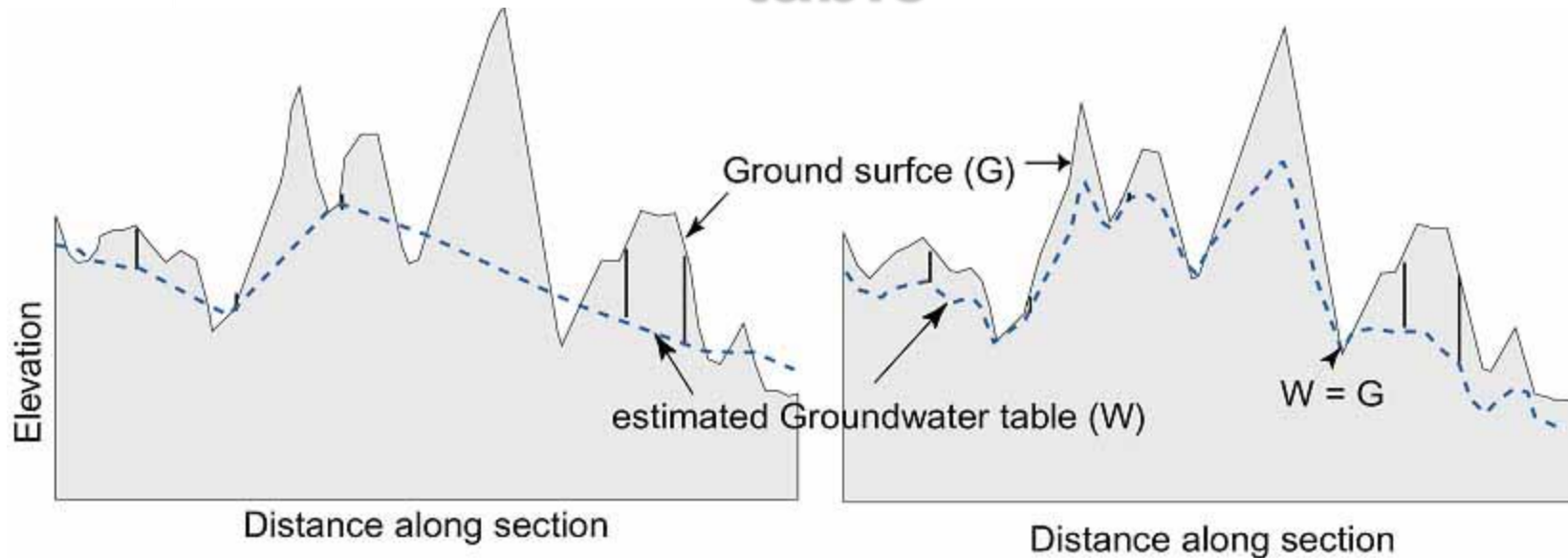




DTW: Depth to Groundwater; $DTW_1 > DTW_2$

- **The groundwater table tends to mimic the overlying topography (except in karst terrain), and is subject to seasonal and decadal variations, depending on precipitation and withdrawal**

GIS software tends to oversimplify natural undulations of the ground water table

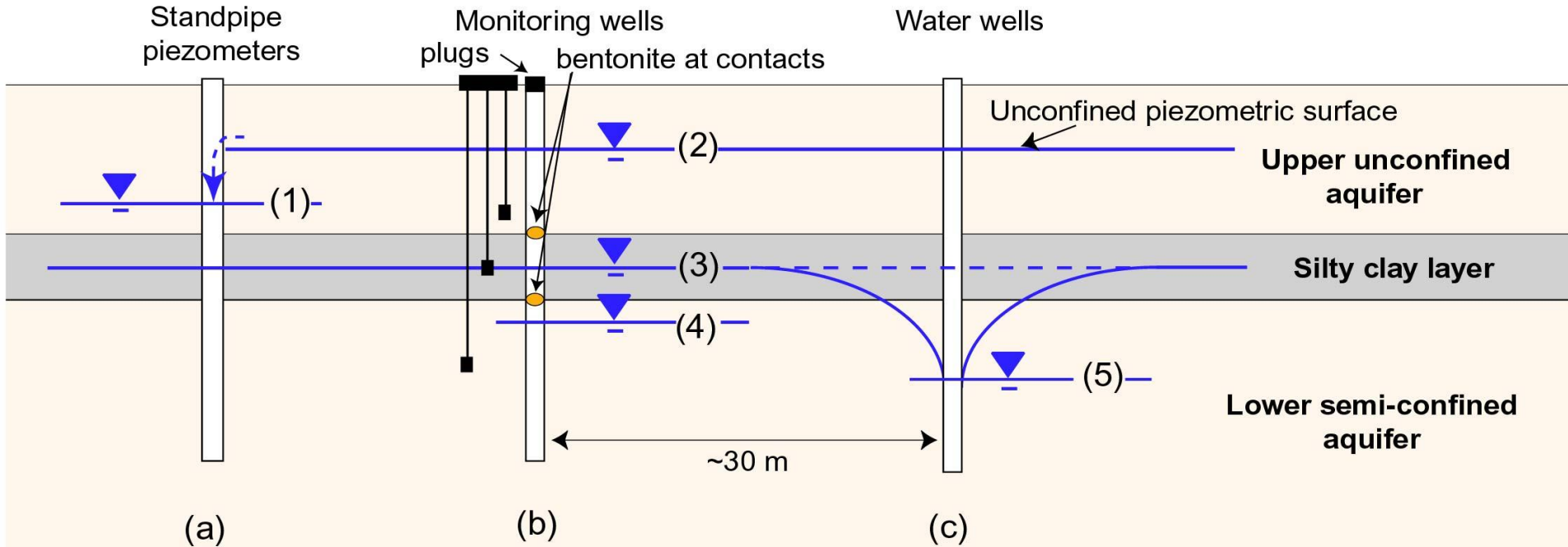


- Predicted ground water table (gwt) using kriging in ArcGIS
- Very unrealistic

- Predicted gwt using cokriging with digital elevation model
- Much more realistic

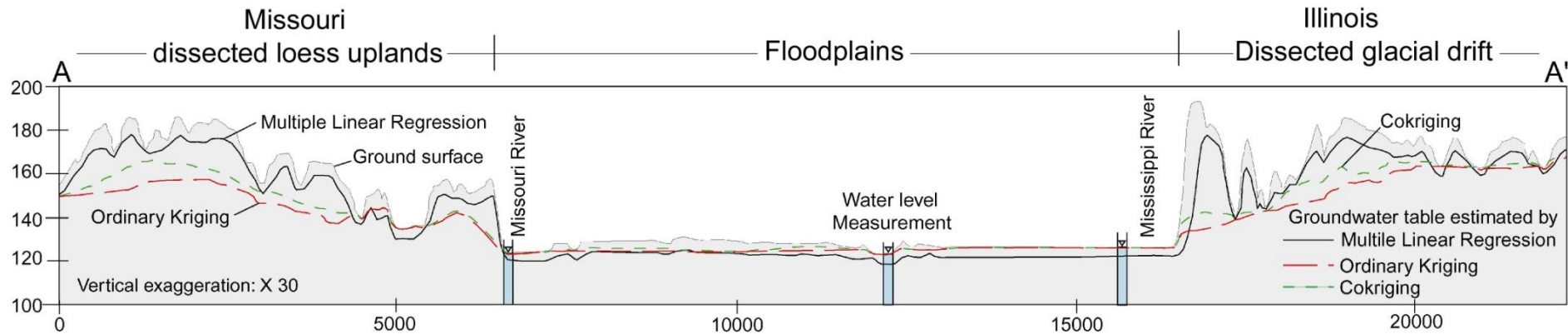
Uncertainty of Reported Groundwater Levels

Measurement errors very common



- (1) - the observed level in a simple standpipe records to combined inflow from the upper unconfined layer and the lower semi-confined layer.
- (2) - the true groundwater level in the upper unconfined aquifer.
- (3) - the true groundwater level of the lower semi-confined aquifer, within the silty clay aquiclude.
- (4) - where groundwater level would typically be recorded, at the time of drilling (atod).
- (5) - the reported groundwater level in an active water well, subject to conical drawdown.

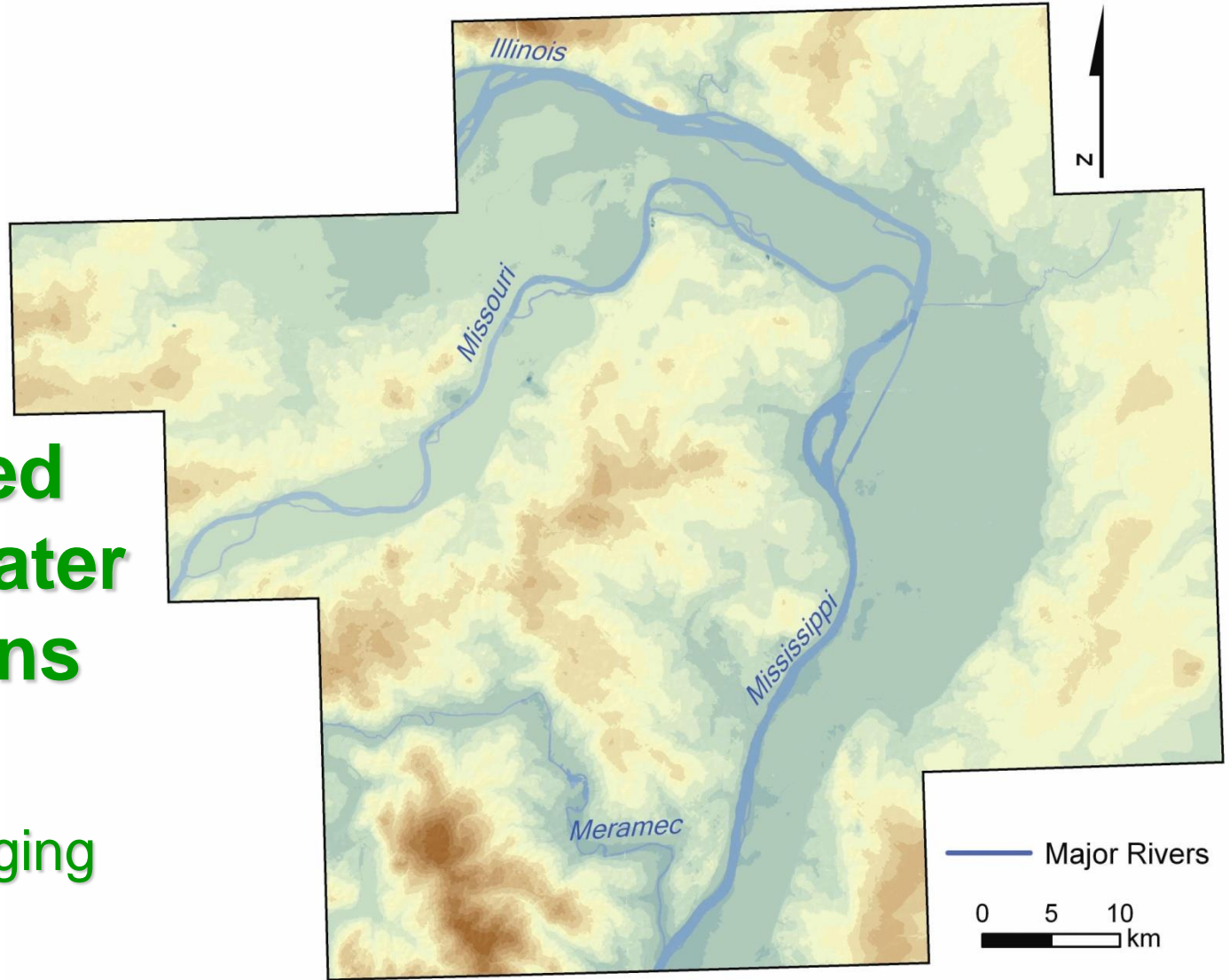
Beware of GIS-driven predictions of groundwater levels



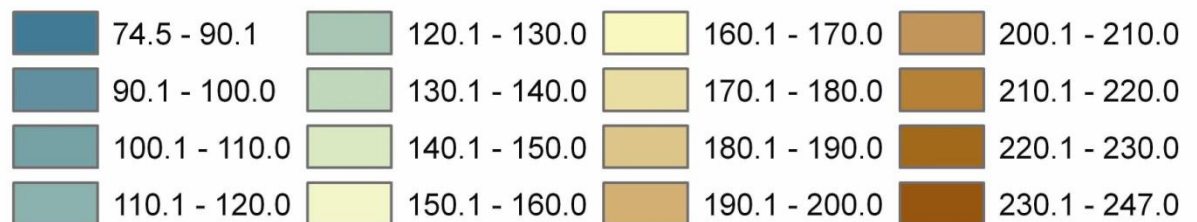
- Most geoenvironmental and geohydrology studies presently employ the geostatistical method of kriging to estimate groundwater levels, using available well records.
- Cokriging that includes dissected ground surface as a covariable usually provides a more realistic prediction.
- Note that multiple regression overestimates in the uplands and underestimates the true values beneath the flood plains.

Predicted Groundwater Elevations

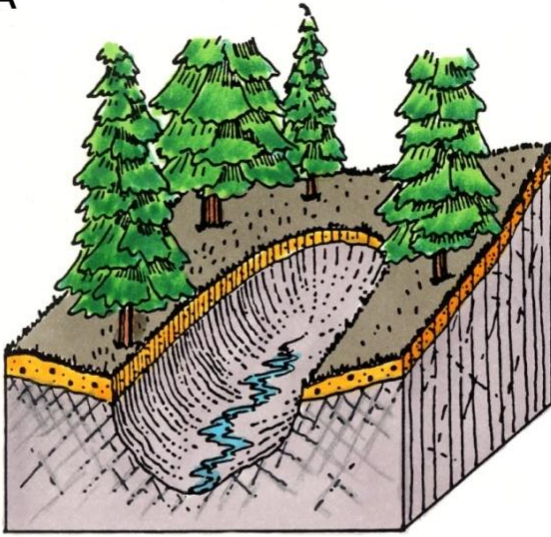
- Produced using cokriging



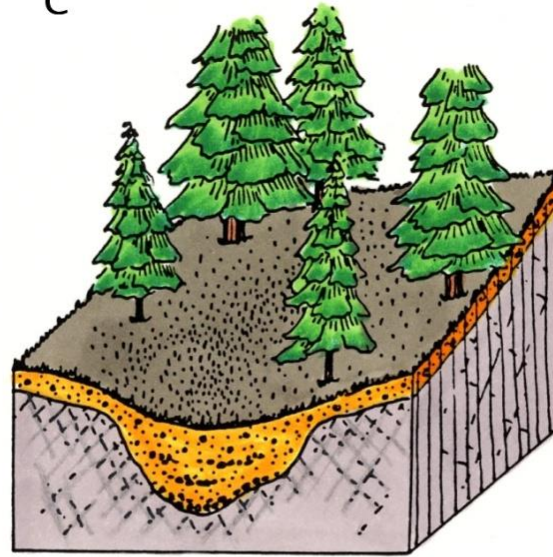
Elevation of Groundwater Table (m)



A

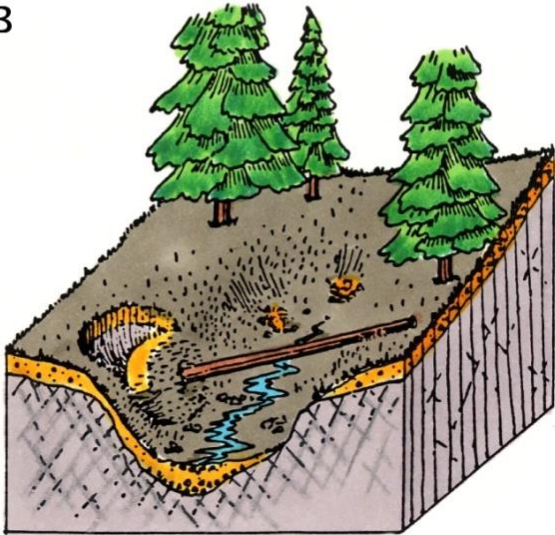


C



The most
common
geomorphic
feature in
America

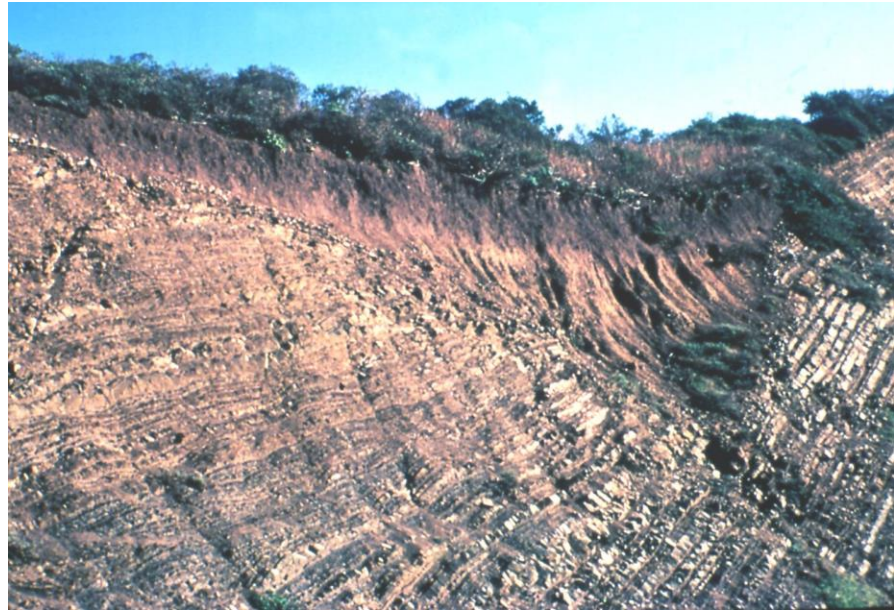
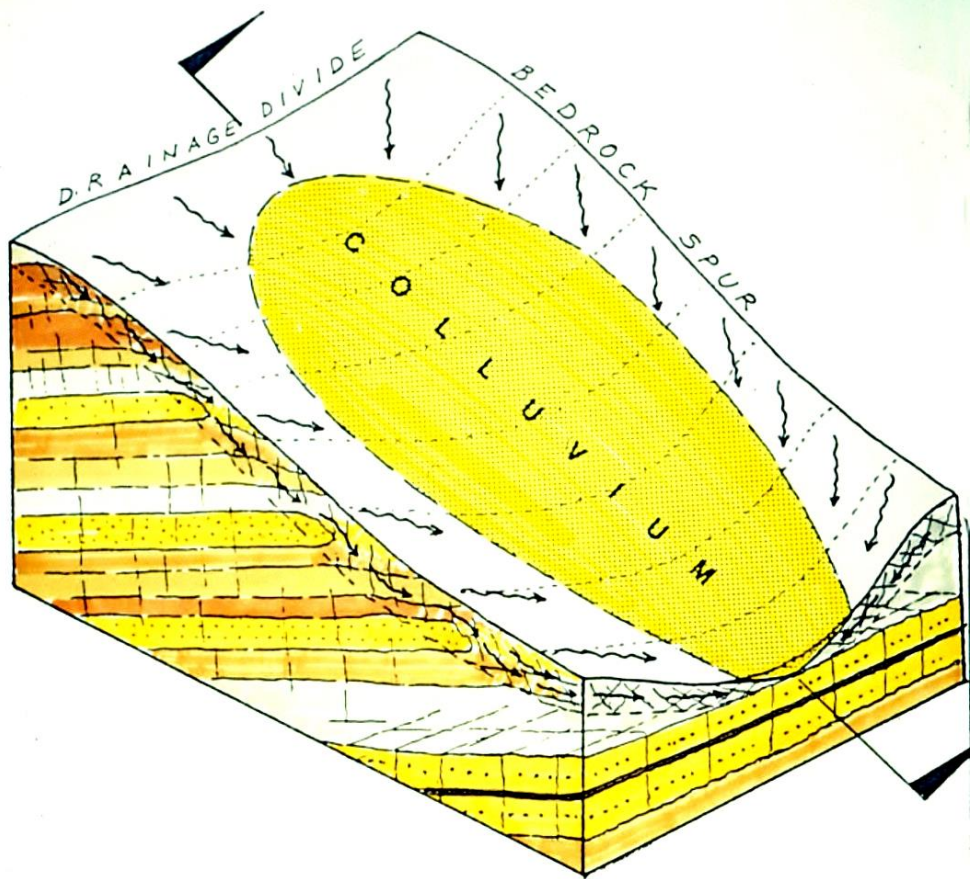
B



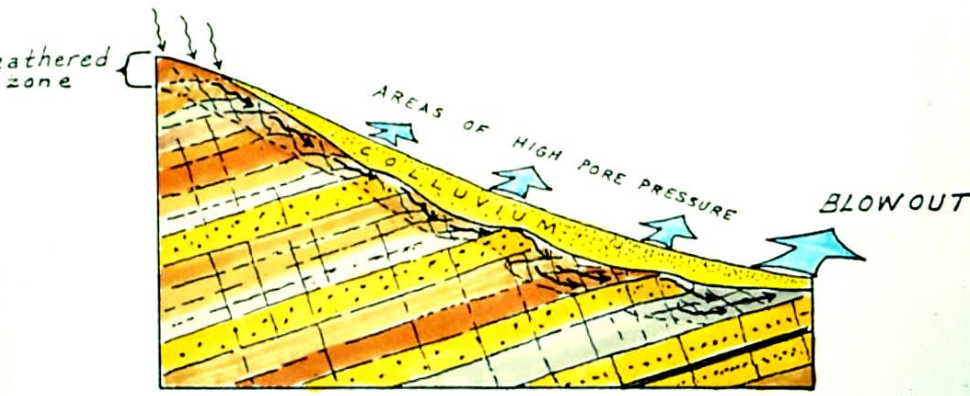
- Evolution of a typical **colluvial-filled bedrock ravine**
- Bedrock hollows are excavated during periods of accelerated erosion, then slowly infill during periods of aggradation



- Colluvial filled bedrock ravines are the most common landform in the world. Shallow groundwater tends to be concentrated towards the mouths of these **zero-order basins**



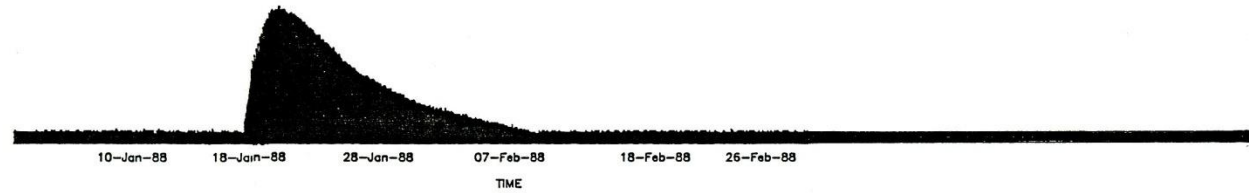
- Most colluvial filled ravines tend to be spoon shaped, as shown here
- **Hydraulic uplift** promoted by seepage in the weathered bedrock tends to concentrate at bedrock “highs”, as shown at lower left





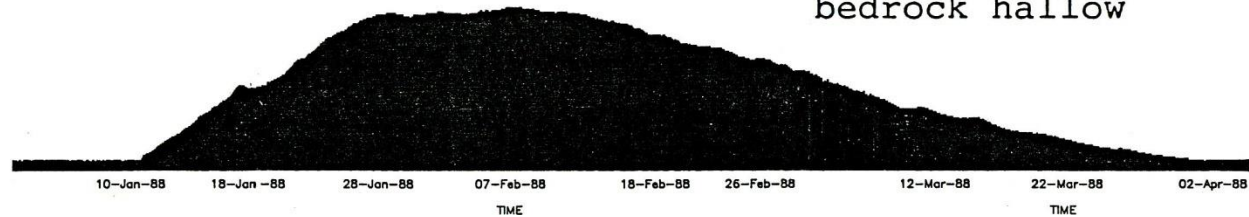
PIEZOMETER DATA
HOLE 8, JAN. 18 – MARCH 2, 1988

PIEZOMETER in
weathered bedrock



PIEZOMETER DATA
HOLE 9, JAN. 18 – MARCH 2, 1988

PIEZOMETER in
colluvial soil
cover of a
bedrock hollow

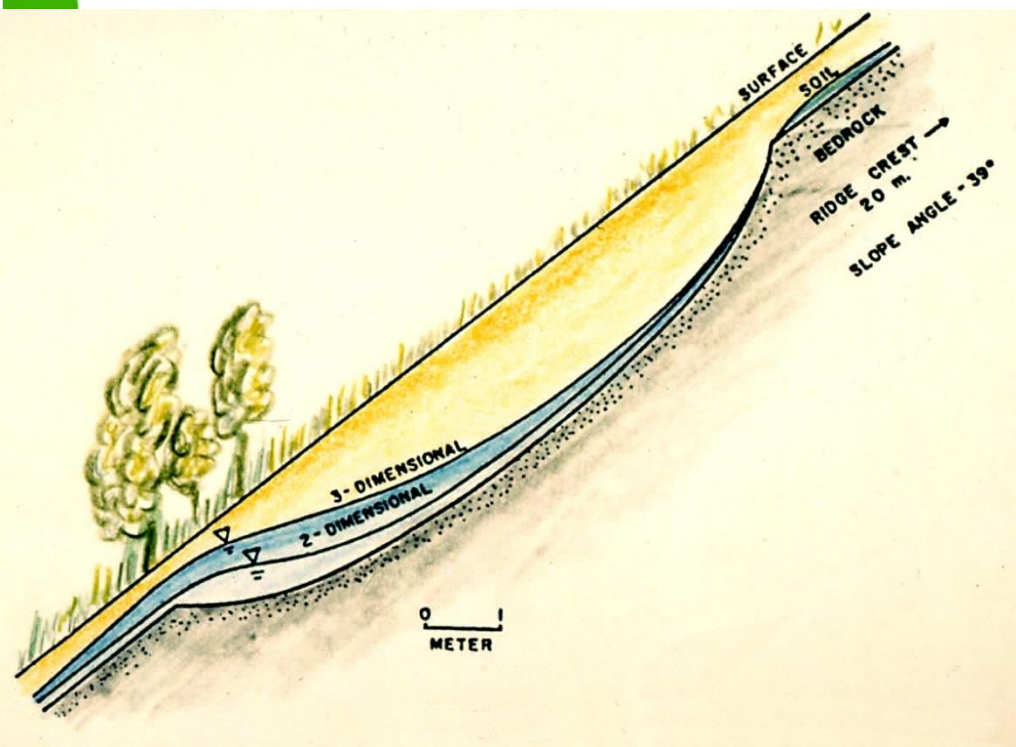


- Water levels recorded by piezometers placed on the bedrock rib adjacent to a colluvial filled ravine (upper diagram), and another in the axis of the colluvial deposit (lower diagram). Note how quickly the water enters the weathered rock and how slowly it drains from the colluvium



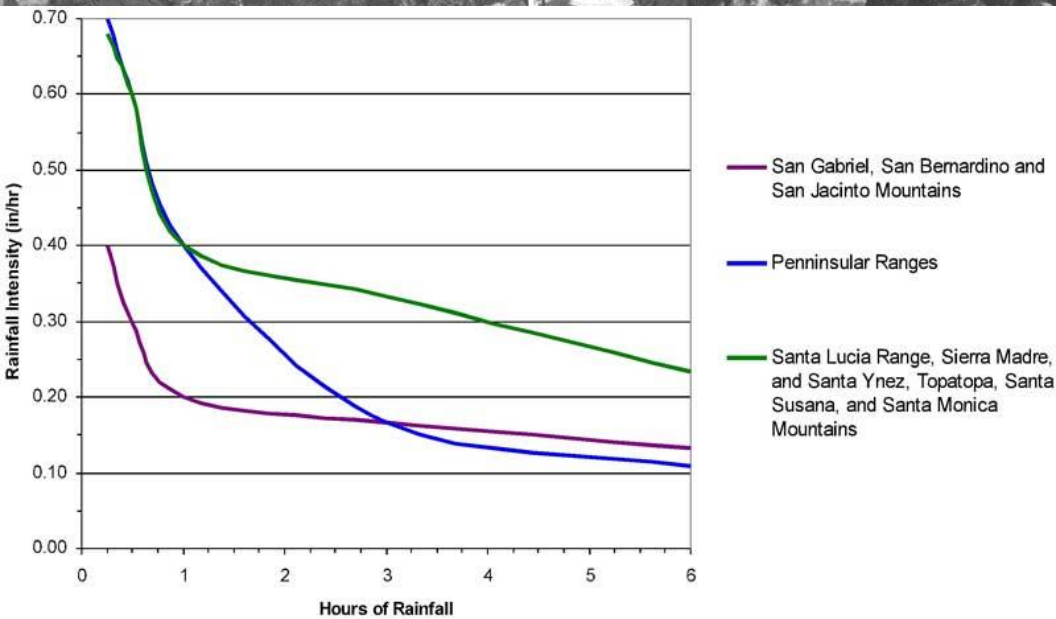
Modeling pore pressures using finite elements

- **Seepage forces** tend to concentrate towards the lower end of a spoon-shaped bedrock ravine, as shown here
- This diagram shows the predicted position of trapped groundwater using 2D and 3D solutions
- Two feet of seepage head can “lift” 4.5 feet of saturated colluvium off the slope



Trigger thresholds for debris flows

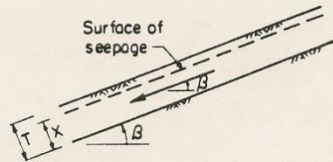
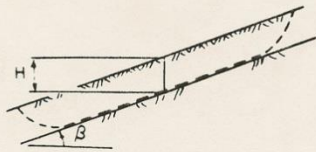
- The USGS has empirically derived rainfall intensity vs duration trigger thresholds for many areas of the USA
- Those at lower left are for Central and Southern California Coast Ranges



Transient Pore Pressures

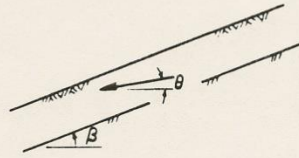


- Transient pore pressures develop during periods of sustained intense precipitation, occasionally triggering explosive “blow-out” failures, like that shown here



Seepage parallel to slope

$$r_u = \frac{x}{T} \frac{\gamma_w}{\gamma} \cos^2 \beta$$



Seepage emerging from slope

$$r_u = \frac{\gamma_w}{\gamma} \frac{1}{1 + \tan \beta \tan \theta}$$

γ = total unit weight of soil

γ_w = unit weight of water

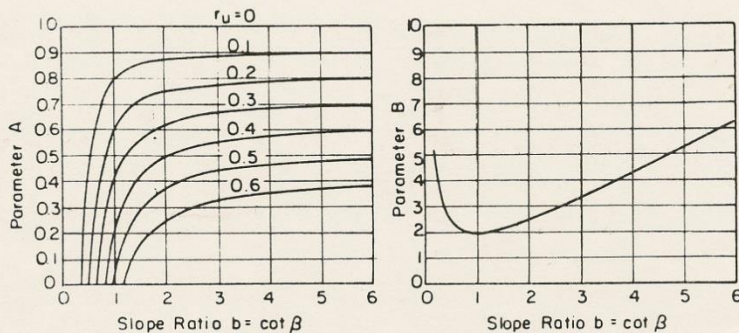
c' = cohesion intercept } Effective Stress
 ϕ' = friction angle }

r_u = pore pressure ratio = $\frac{u}{\gamma H}$

u = pore pressure at depth H

Steps:

- ① Determine r_u from measured pore pressures or formulas at right
- ② Determine A and B from charts below
- ③ Calculate $F = A \frac{\tan \phi'}{\tan \beta} + B \frac{c'}{\gamma H}$



- An **Infinite Slope Analysis** can be performed to illustrate how a few feet of excess pore pressure can trigger such failures

BASIC PRECEPTS OF GROUNDWATER FLOW

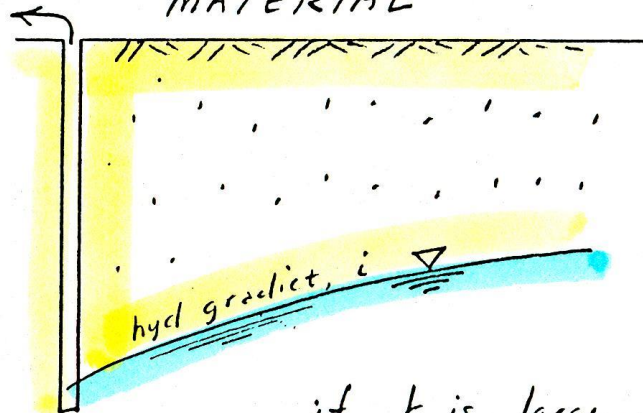
Darcy's Law:

$$Q = kiA \quad ;$$

or Quantity of Flow = (permeability) (hydraulic gradient) (Cross Sectional Area of Flow)

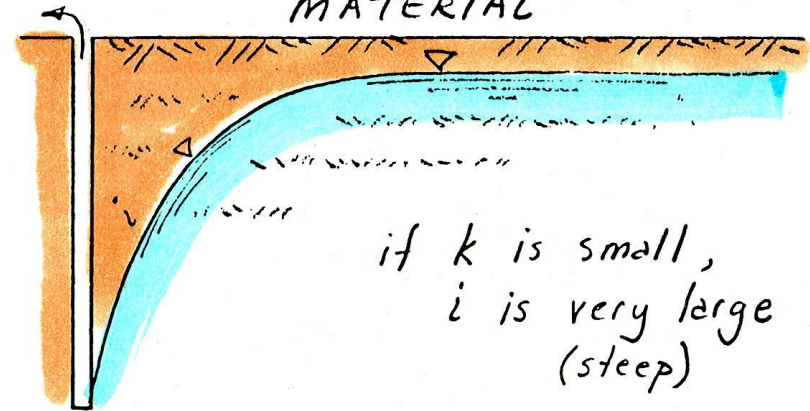
e.g. Well drawdown

IN SANDY, PERVIOUS MATERIAL



if k is large,
 i is small

IN CLAYEY, LESS PVIOUS MATERIAL



if k is small,
 i is very large
(steep)

- Permeability (k) and the hydraulic gradient (i) are more or less inversely proportional (see sketches above)
- Low permeability materials don't tend to transmit significant quantities of water