LESSONS ON SITE CHARACTERIZATION GLEANED FROM FORENSIC STUDIES

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KARL TERZAGHI

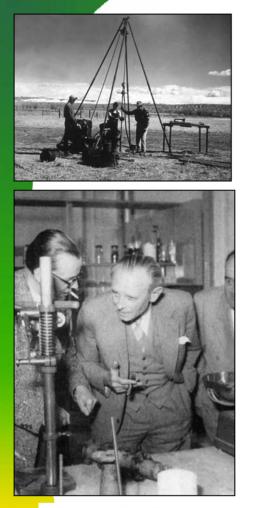
- Austrian engineer 1883-1963
- Graduate training in geomorphology
 - World traveler and astute observer of nature
- Fathered soil mechanics
 - Awarded 4 Norman Medals
- Promoted appreciation of engineering geology by civil engineering profession



TERZAGHI'S METHOD

- Study geology and geomorphology of region surrounding project site
- Gather all forms of existing data, including geologic, soils, hydrologic, meteorlogic
- Identify "missing gaps" in geo information
- Make on-the-ground reconnaissance of the site; note dominant erosional processes
- Formulate a working hypothesis regarding the nature of likely subsurface conditions
- Develop plan of site exploration, designed to fill missing gaps





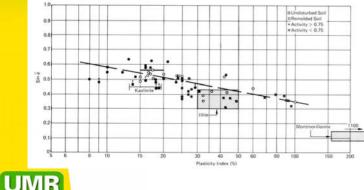
TERZAGHI'S REPORTS

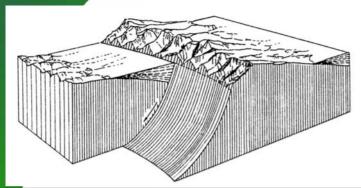
His first report was a summary of site geology and what he perceived to be the data gaps

The second report summarized the site specific investigation and either characterized site conditions or demanded additional work

Terzaghi made thorough use of ongoing construction observations to verify assumed site conditions

> Always made provisions for changing plans whenever different conditions were observed









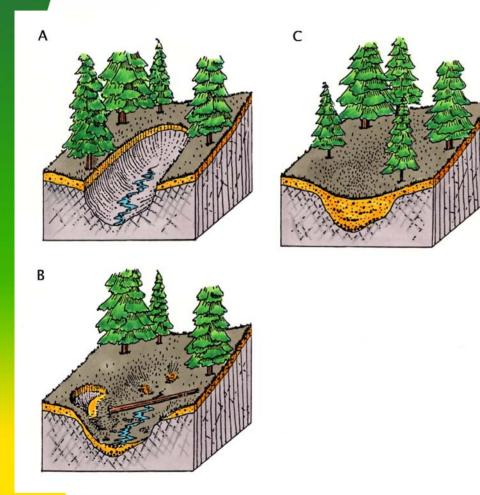
EXPERIENCE

Geologists only recognize those features with which they have accumulated substantive experience

Mis-characterization most often occurs when people making the examinations have little previous experience in the area

Geologists use different technical terms, based on their experience, e.g. listric faults versus landslide slip surfaces, or colluvium vs slide debris

THE COLLUVIUM CATCH ALL



 Colluvium is loose, heterogeneous, and incoherent mass of soil material and/or rock fragments deposited by rainwash, sheet wash, or downslope creep

Block diagrams showing cycle of colluviual deposition and erosion

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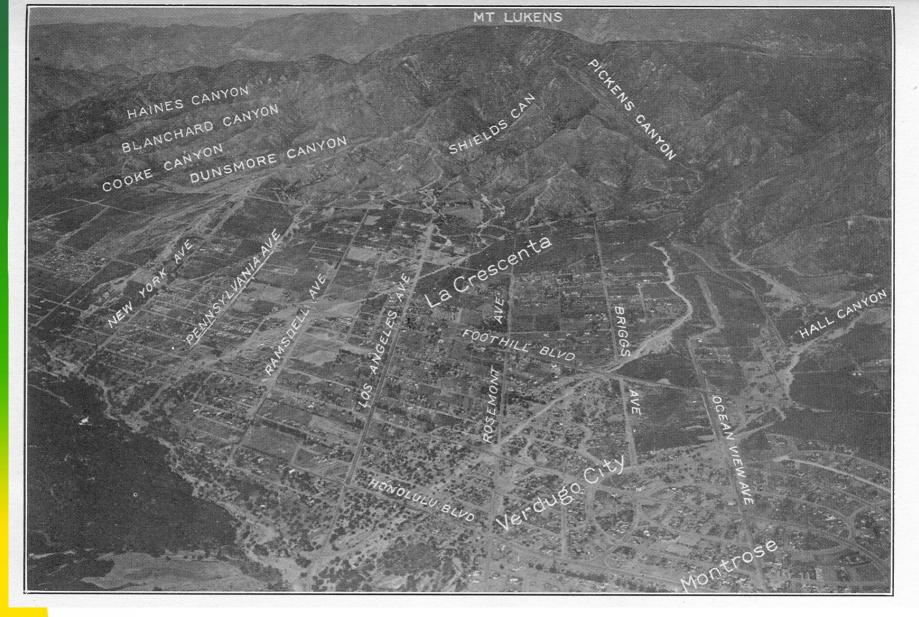
Technical terms connote genesis

Slide debris is often mistaken for colluvium because the geologist doesn't see a slip surface

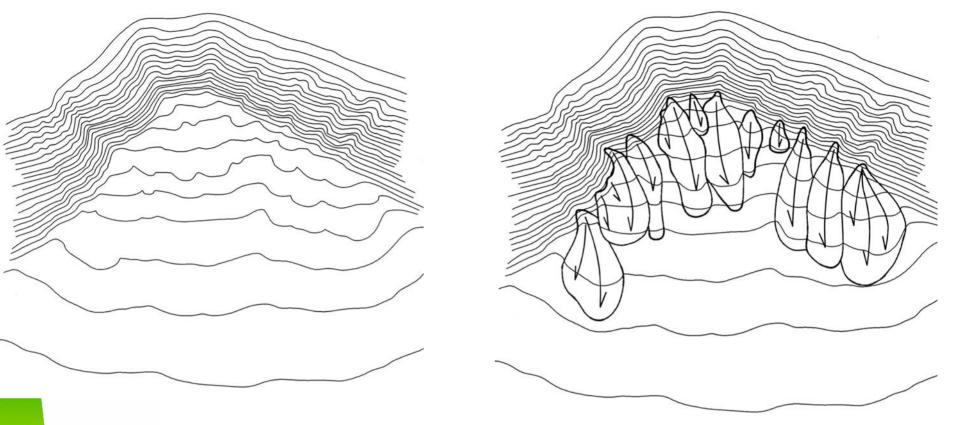


Loose soil debris on a grassy slope will be deposited without a shear surface





The Dec 31, 1933-Jan 1, 1934 Montrose-LaCresenta debris flows damaged or destroyed 600 homes and killed 44. 600,000 cubic yards of material was deposited on the fan in one evening.



Many mapped alluvial fans (Qal) are actually debris fans, constructed by countless series of coalescing debris flows. One trench can expose the character and genesis of underlying material. If it is matrix-supported, it was likely deposited in a debris flow.



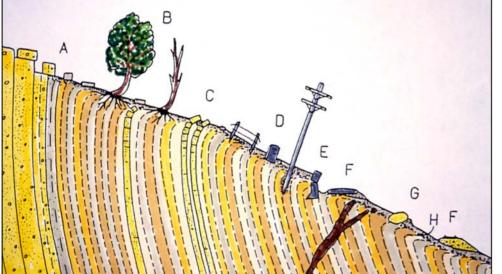




MATRIX SUPORTED MATERIALS

- Debris flow deposits are typically deposited on channel slopes around 10%
- They are characterized by fine grained matrix between clasts, large variation in clast and particle sizes, and often, by inverse sorting



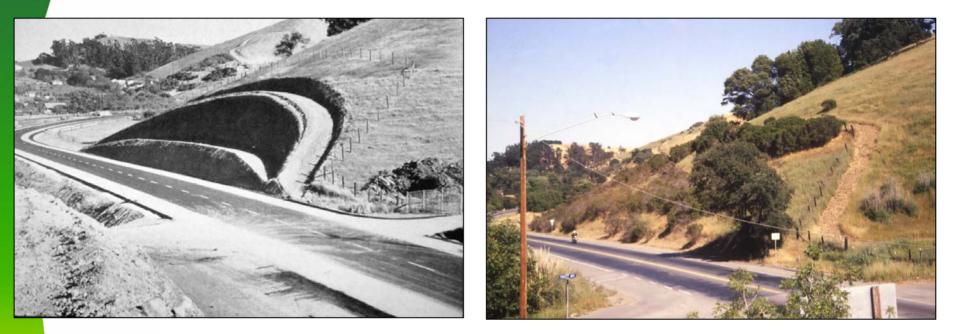


SLOPE CREEP

- Geologists are often asked to estimate depth and magnitude of slope creep
 - Creep is a function of soil plasticity, slope inclination, seasonal moisture variation and height of slope

Slope creep can exert passive soil pressures against embedded foundations

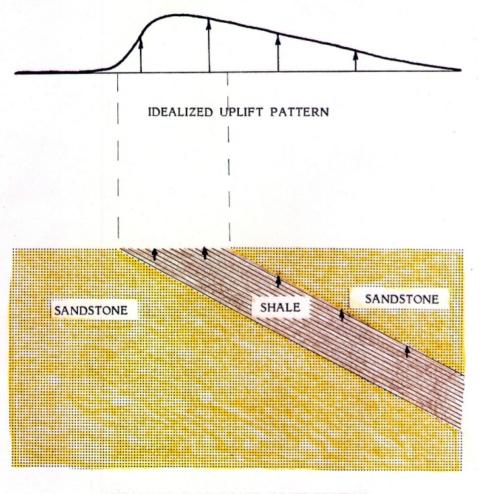
AGING FACTORS



- Clayey soils and argillaceous materials are susceptible to slope creep. Be careful to advise clients of likely movements and need for ongoing maintenance
- These comparative views of the same highway cut were taken in 1954 and 1986, 32 years apart. The benches were gone.



DIFFERENTIAL HEAVE



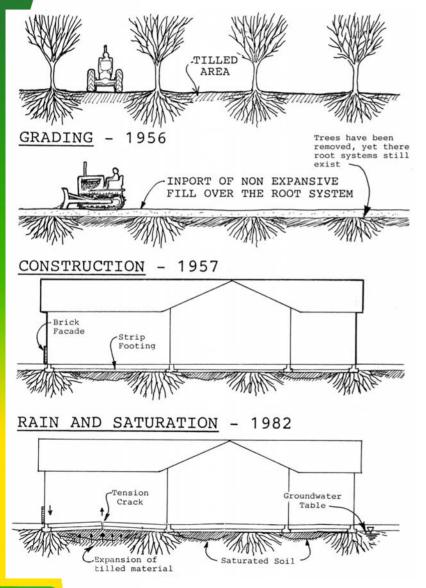
IDEALIZED SUBSURFACE CROSS-SECTION

Differential heave is a recurring problem whenever expansive clay shales are sandwiched between non-expansive beds, as sketched at left

 Expansive soils are the #2 cause of property loss in the United States, second only to dry rot



PAST LAND USEAGE



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 A preliminary check of aerial photos is essential to understand what a site was used for prior to anticipated construction

 This shows the case of a subdivision in San Jose, CA which was built over an old orchard



FILLS AND OLD CHANNELS

Check old maps and air photos to see where channels and sloughs were located prior to infilling, draining or redirection. This is a common problem in urban areas



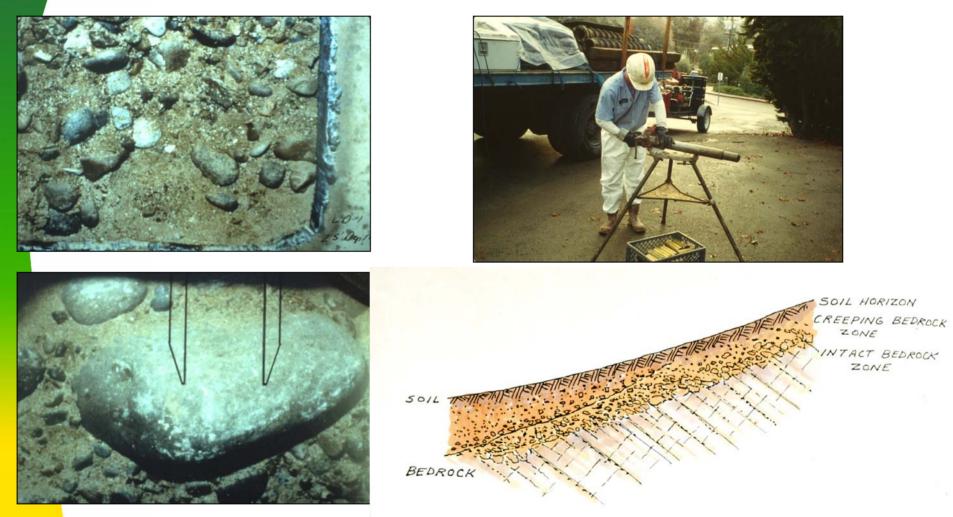
INTERPRETING SUBSURFACE DATA



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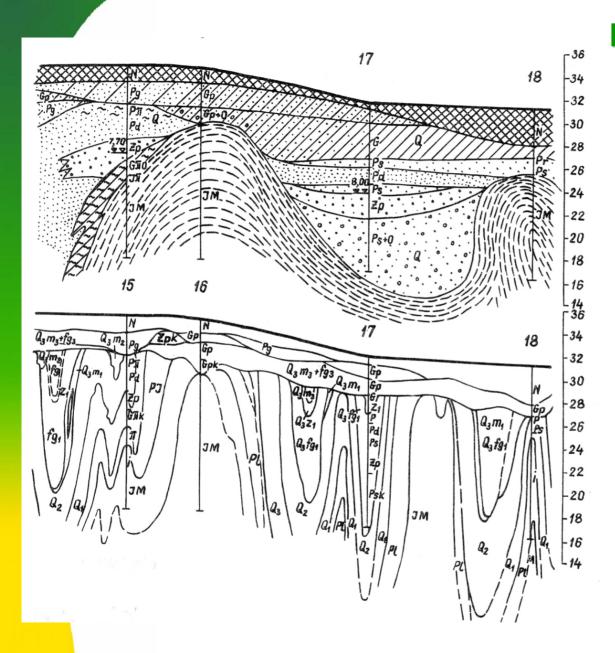
- Soils engineers tend to focus on providing foundation recommendations, not on detecting geologic contacts
 - Proper interpretation of subsurface information requires experience on part of the drillers and the logging geologist
- Circulation problems can negate accurate assessments of cuttings depth and weak seams of horizons are easily missed

CLASTS LARGER THAN SAMPLER SHOE



Bedrock contacts are commonly misidentified when the sampler shoe encounters a floater of greater diameter, as sketched at lower left





MISINTERPRETATION OF DATA

Subsurface data is easily misinterpreted when working in tectonically active terrain.

This shows before construction (upper) and after excavation (lower) cross sections of the same site

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Exploratory trenches can effectively expose geologic structures and the nature or genesis of key contacts



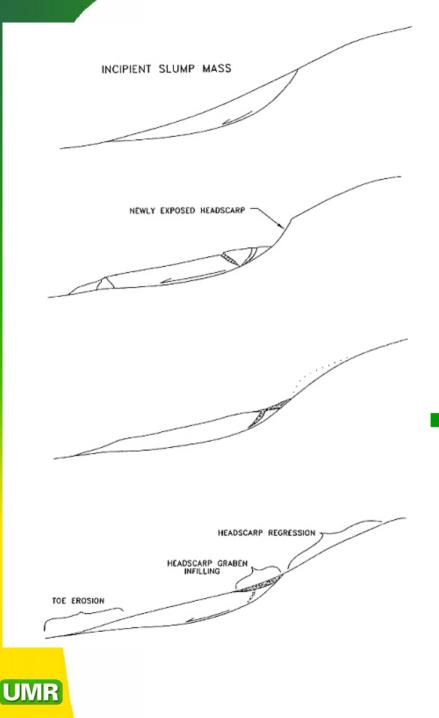




RECOGNIZING OLD LANDSLIDES

- Old landslide features can be difficult to discern, even for experienced engineering geologists
- One needs to be "looking" for such features in order to recognize them
- Case at lower left was old slide reactivated on slip plane inclined just 3 degrees, without any visible moisture



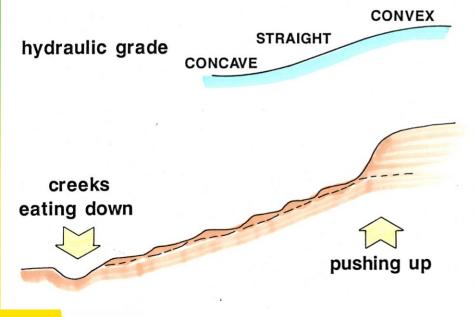


TOPOGRAPHIC EXPRESSION



A key indicator of past landslippage is anomalous topographic expression, which becomes increasingly mollified with age, as shown in profiles at left





TOPOGRAPHIC EXPRESSION OF LANDSLIDES

Landslides tend to form coalescing complexes, with hummocky topography and deranged drainage patterns. Toe undercutting or tectonic uplift are common triggers





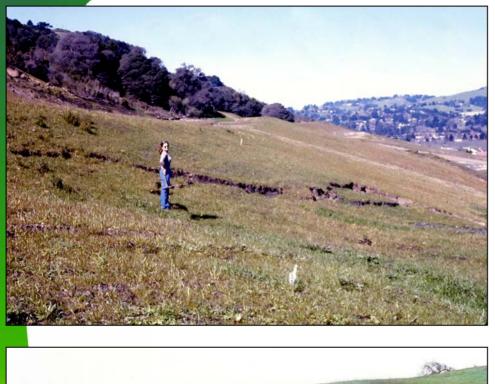
RECOGNIZING OLD LANDSLIDES





Recent landslides like those at left are easily noticed, but older complexes can be difficult to discern, especially under thick tree cover. The hummocky topography in the right image is typical of landslide-prone terrain.







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INCIPIENT LANDSLIPPAGE

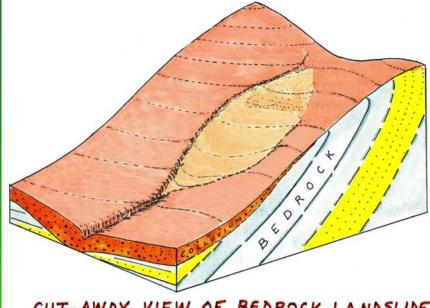
Incipient landslides are those that are beginning to fail, but have not fully ruptured They are easily identified by their arcuate tensile scarps crossing otherwise intact slopes

DEEP- SEATED BEDROCK LANDSLIDE

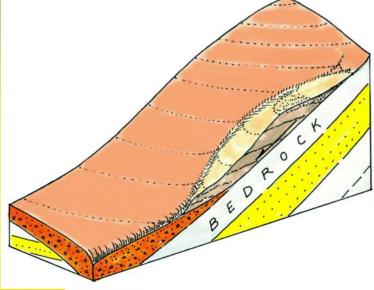
OLD BEDROCK LANDSLIDES



Old bedrock landslides can be very difficult to identify. Key indicators are anomalous isolated benches and converging parallel drainages



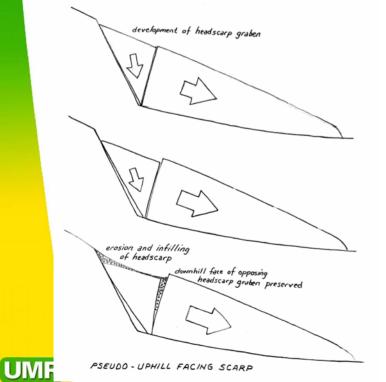
CUT-AWAY VIEW OF BEDROCK LANDSLIDE



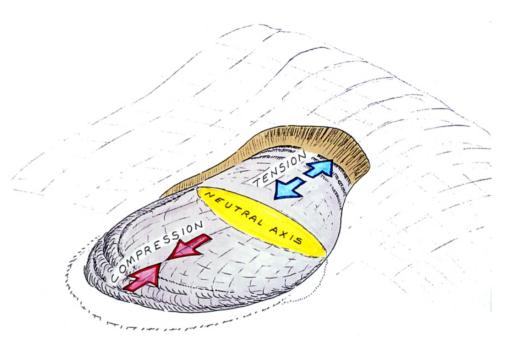
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SIMPLE TRANSLATIONAL BLOCK GLIDE LANDSLIDE



HEADSCARPS



The arcuate nature of a landslide headscarp is due to tensile pull-apart as the slide mass translates downslope. It is easy to estimate the depth of sliding from the opposing scarps

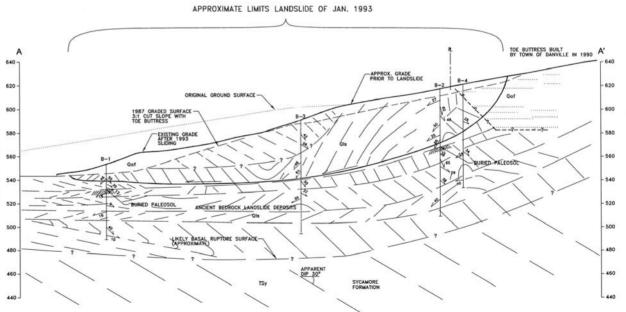


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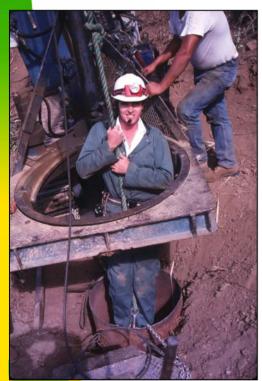
LEVATION

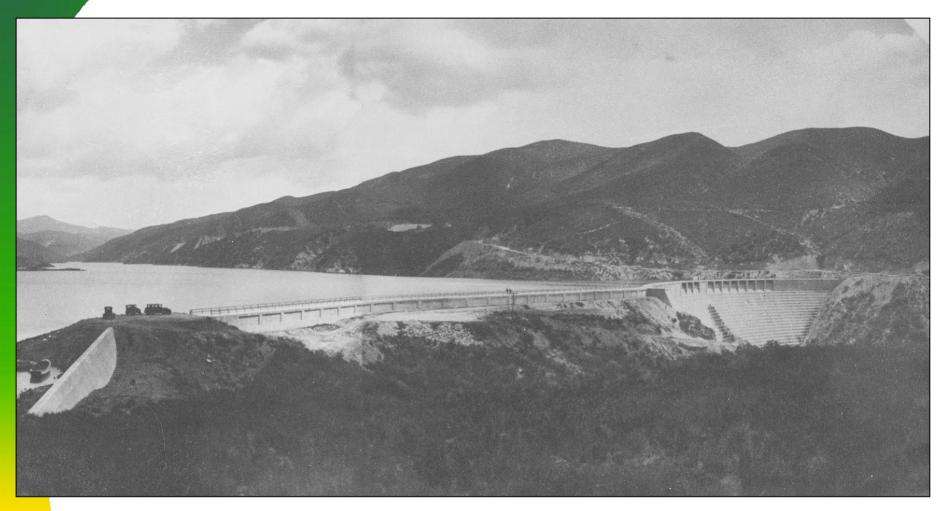
BUCKET AUGERS

Bedrock landslides are best explored using large diameter bucket auger borings, which can be downhole logged



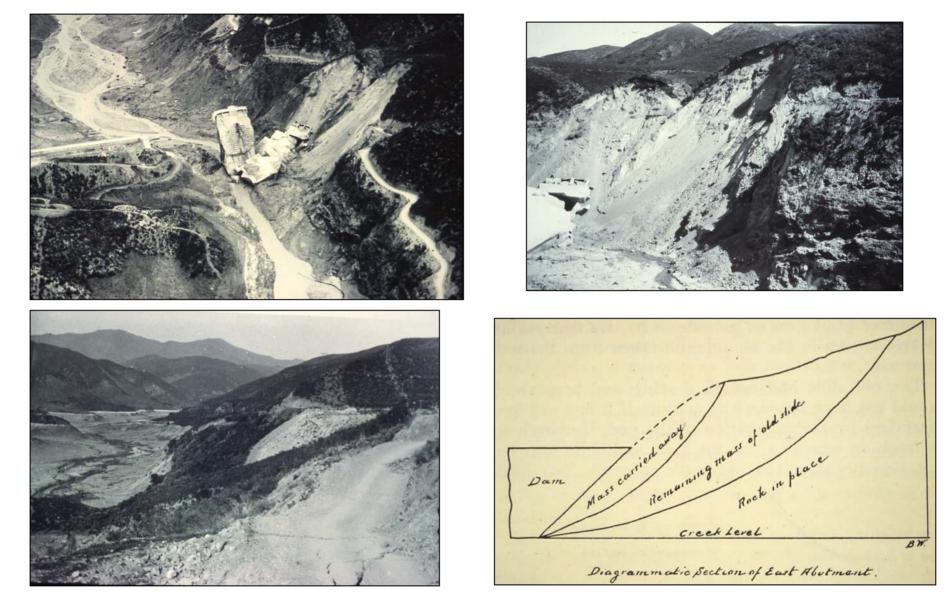
GEOLOGIC CROSS-SECTION A-A' SHOWING APPROXIMATE STRUCTURE WITHIN LANDSLIDE MASS



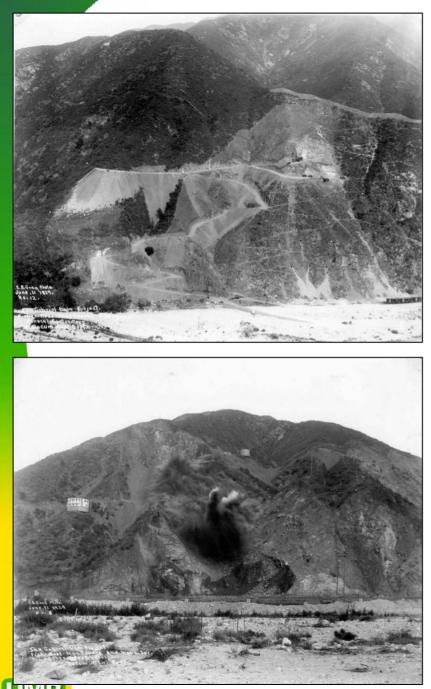


St. Francis Dam was unknowingly constructed against a paleolandslide developed in pre Cambrian age Pelona Schist in 1924-26.





535 million m³ of schist in the dam's left abutment detached and slid downslope, destroying the dam. Relaxation movements of up to 3 m were noted at elevations more than 60 m above the left abutment

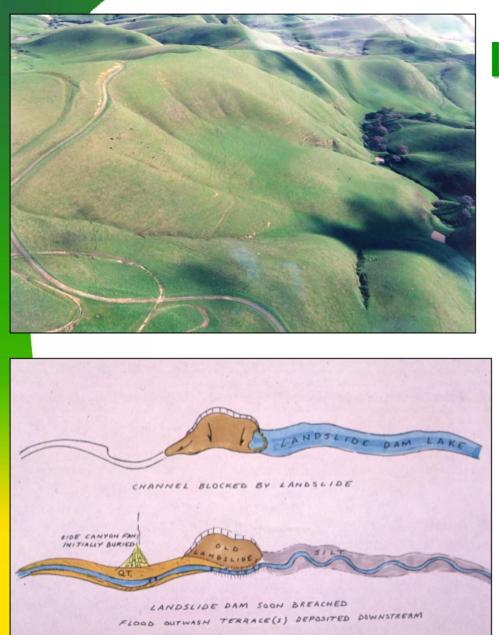


SAN GABRIEL DAM

- The San Gabriel Dam at the Forks Site would have been the largest dam in the world when construction began in September 1928
- These views show right abutment excavations at the Forks site in June 1929
- Plans called for 200,000 yds³ excavation on each abutment



Three views showing before (left) and during (middle) and after (right) the massive detonation of the dam's right abutment. The contractor detonated 87,430 kg of dynamite. On September 16, 1929 a massive slide of the right abutment occurred, bringing down 153,000 m of additional rock. This led to an investigation and eventual cancellation of the project by the State of California.

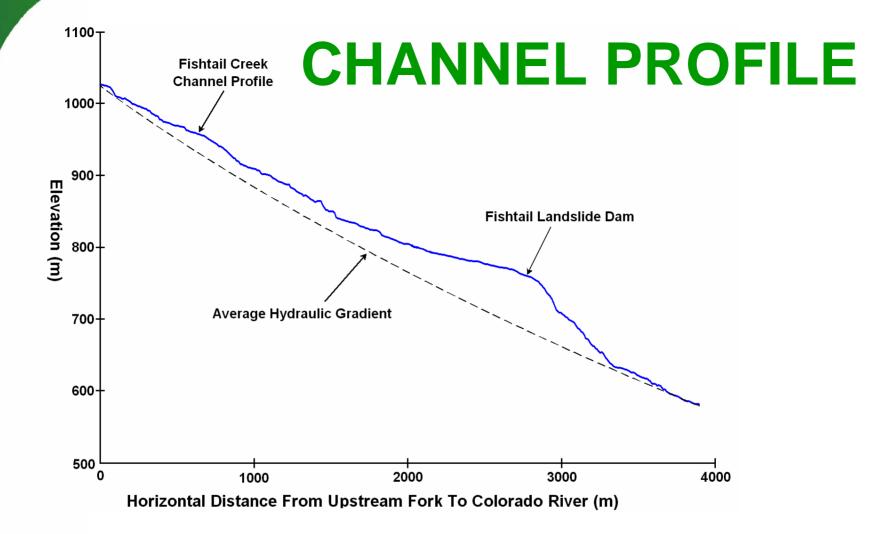


LANDSLIDE DAMS



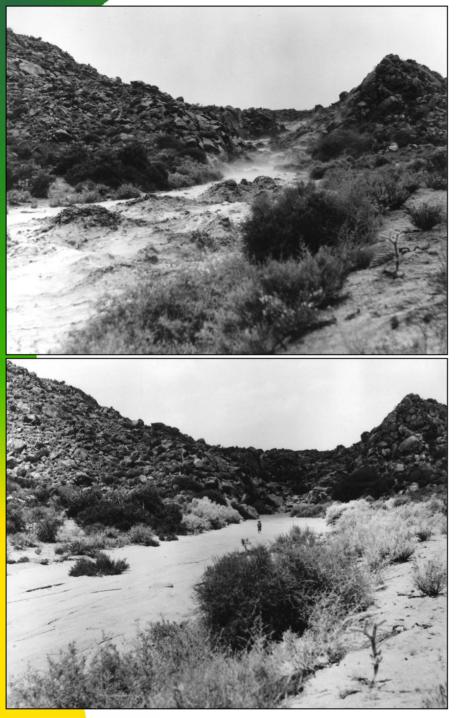
Landslide dams perturb channel profiles. Lessened gradient upstream of blockage and increased gradients through the obstruction, as shown above. Channels often swing around the obstructions in a convex outward pattern.





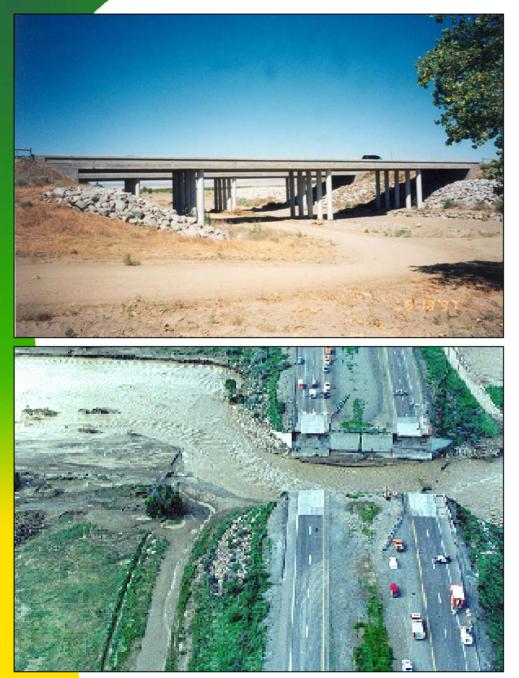
Example of a channel profile perturbed by landslide damming in Fishtail Creek, Grand Canyon. Note lower gradient upstream of obstruction and increased gradient on downstream side.

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CHANNEL SCOUR

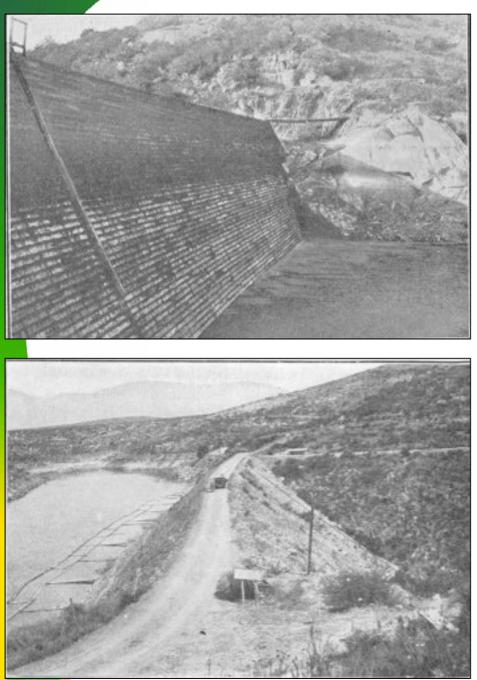
- Engineering geologists are often asked to estimate depth of channel scour, especially for bridges
- These are during and after views of a desert flash flood in June 1969 which developed 8 feet high standing waves. Note children for scale in lower image, taken a few weeks later.



Arroyo Pasajero Wash Out

A flash flood on March 10, 1995 brought a peak flow of 33,000 cfs through a constricted overcrossing for Interstate 5 near Coalinga, CA. The channel was downcut by almost 30 feet, undermining the bridge caissons and dropping the twin spans. 7 people were killed.



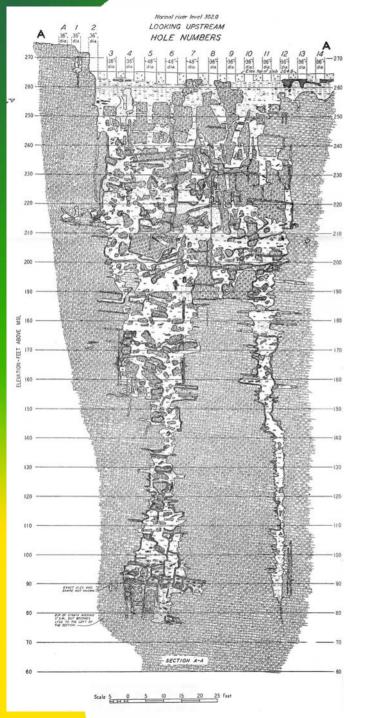


DAMS THAT WOULD NOT HOLD WATER

Escondido (upper) and Morena (lower) Dams were both constructed around 1900 in southern California. They lost large volumes of water through alluvial gravels beneath the embankments.

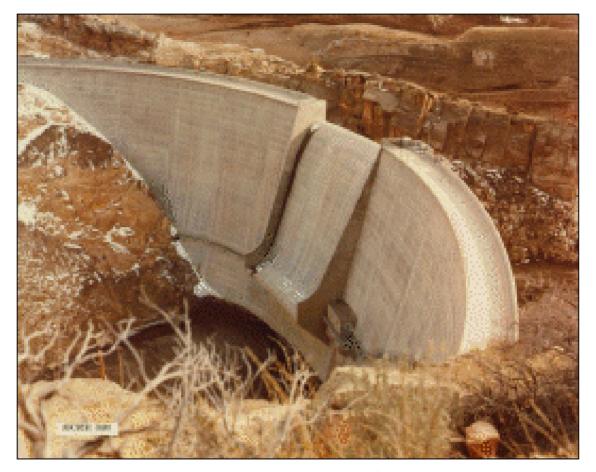
Escondido Dam never filled beyond 2/3 capacity, losing 100,000 gpd. Morena lost between 33,500 to 58,000 gpd, depending on head.

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KARST FOUNDATIONS

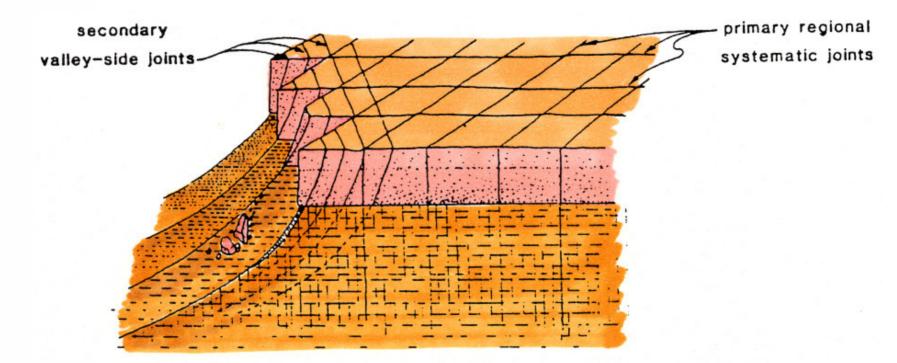
- Cross section through the foundation for Kentucky Dam on the Tennessee River (TVA, 1949).
- The Pleistocene Valley floor was about 70 feet lower than at present, with deep solution cavities extending 200 feet beneath the channel
- 734,000 sacks of cement were pumped into the foundation during construction (20,000 m³)



Anchor Dam was built by the U.S. Bureau of Reclamation 56 km west of Thermopolis, WY in 1957-60. Reservoir water seeps downward into karstified redbeds of the Permian age Goose Egg and Triassic age Chugwater formations. The dam has never retained any significant volume of water.



OUT-OF-PLANE DISCONTINUITIES



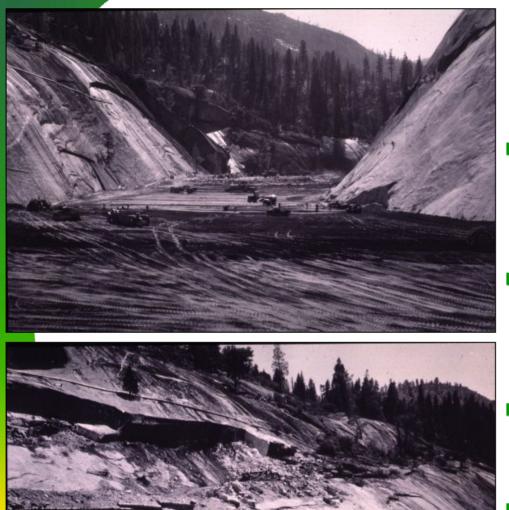
Valley-side stress relief joints tend to form parallel to free cliff faces, in response to stress changes engendered by excavation of the valleys, slope creep and thermally-induced stresses. They are dangerous because they cannot be seen in casual mapping.





Exfoliation joints exposed in glacial cirque of Little Shuteye Pass, in California's Sierra Nevada Mountains, in the Mt. Givens granodiorite



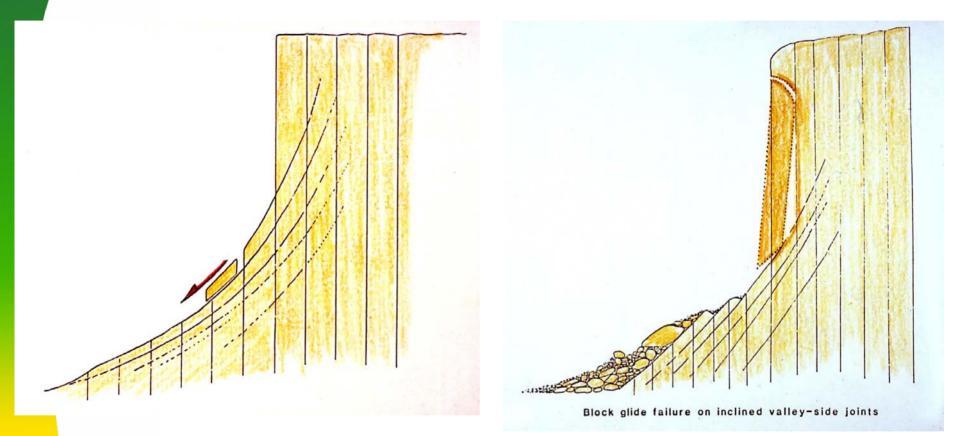


SHEET JOINTS

- Sheet joints, valley side joints or exfoliation joints all describe the same features
 - When the valley side is excavated new joints form beneath the unloading
- This caused problems at Mammoth Pool dam site in California, shown here
- Terzaghi recommended placing fill for dam and grouting the joints as construction progressed



VALLEY-SIDE JOINT FAILURE MODES



Valley-side joints are particularly treacherous because they are usually inclined at close to 45-Ø/2 degrees from vertical (around 60 degrees), which offers the least shear resistance to slippage.

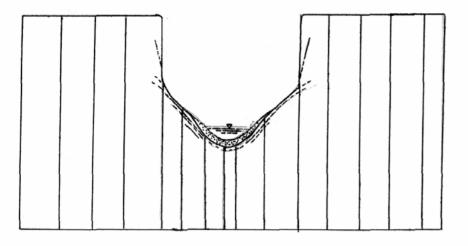




Catastrophic rockfall typical of steeply inclined bluffs with out-of-plane discontinuities. This example is from 6 Mile Wash in Marble Canyon, Arizona



HAUNCHES OF GLEN CANYON DAM



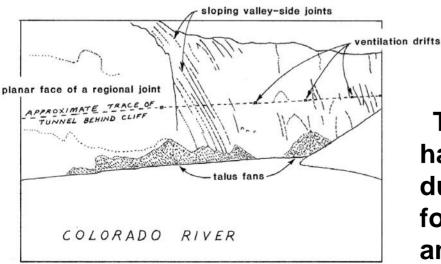


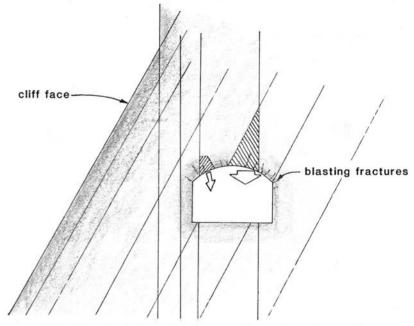
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Glen Canyon damsite has prominent haunches formed by curvalinear valley side joints. The upper vertical walls are controlled by regional systematic joints.

GLEN CANYON TUNNEL FAILURE



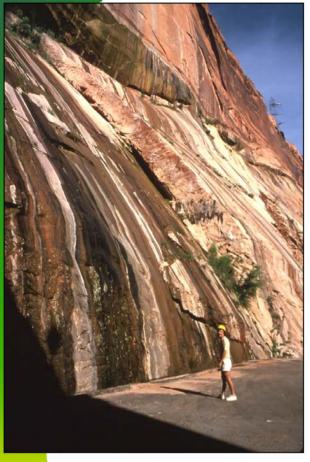




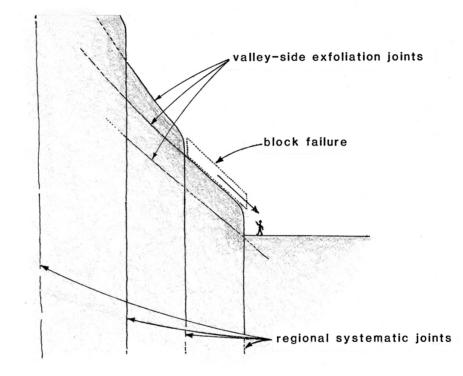
Roof cave-ins in the Glen Canyon Powerplant Service Tunnel during construction in 1958.

The Glen Canyon Powerplant Tunnel had a series of deadly block failures during excavation because of wedges formed between inclined valley side and systematic joints, sketched above





BLOCK FAILURES



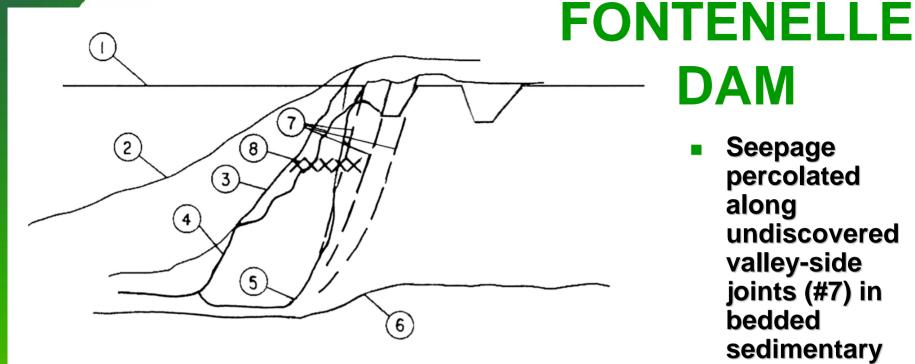


Valley side joints are inclined. Where these intercept near vertical systematic joints, massive blocks are formed on steep inclines. When these are undercut the blocks slide off. A little bit of water hastens the process by decreasing the friction markedly.



Fontenelle Dam came perilously close to failing catastrophically in Sept 1965 during its initial filling. Seepage emanated from the right abutment. The same failure mode befell Teton Dam 11 years later.





Right Abutment Detail.

- (2) Downstream abutment slope.
- (3) Abutment slope at dam centerline.
- (4) Bedrock surface at dam centerline.
- (5) Upstream abutment slope.
- (6) Lower limit of grout acceptance.
- (7) Open cracks.
 - (8) Approximate location of leak.

DAM

Seepage percolated along undiscovered valley-side joints (#7) in bedded sedimentary rocks at position #8 8 lines of grout holes were then drilled and filled with 203,500 sacks of cement

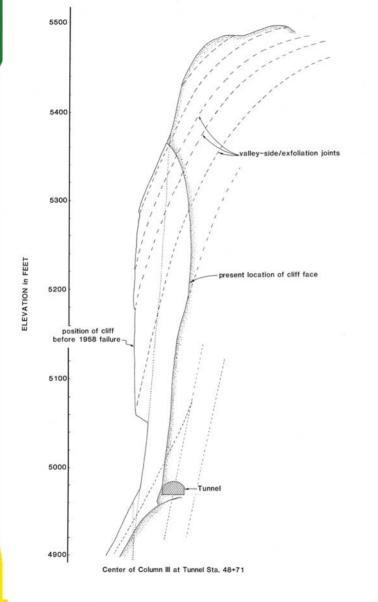
Plan :

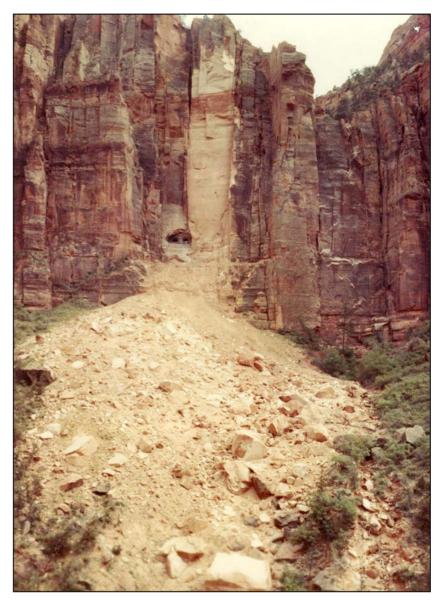
- (A) Embankment.
- (B) Spillway.
- (C) Canal outlet.
- (D) Road.
- (E) September 3, 1965, leakage and slough.
- (F) Open cracks.

Section :

(1) Dam crest.

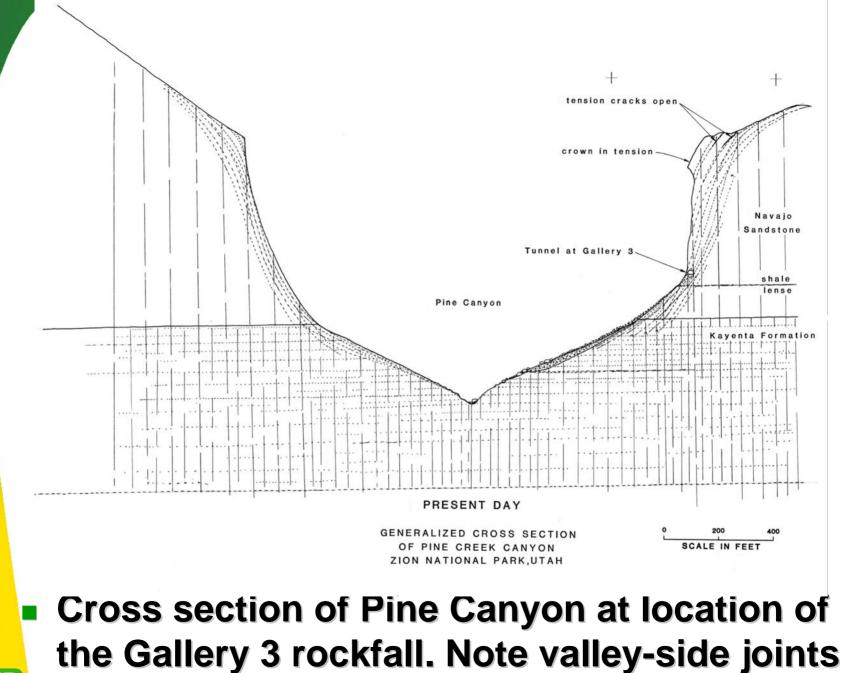




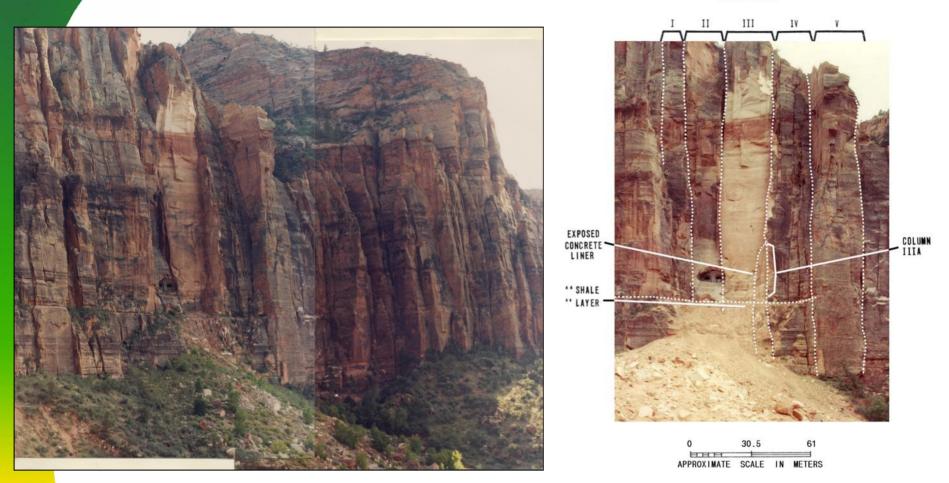


The Zion-Mt Carmel Tunnel Gallery 3 failure occurred in April 18, 1958, spilling 84,000 tons of Navajo Sandstone onto the slope below and damaging the tunnel

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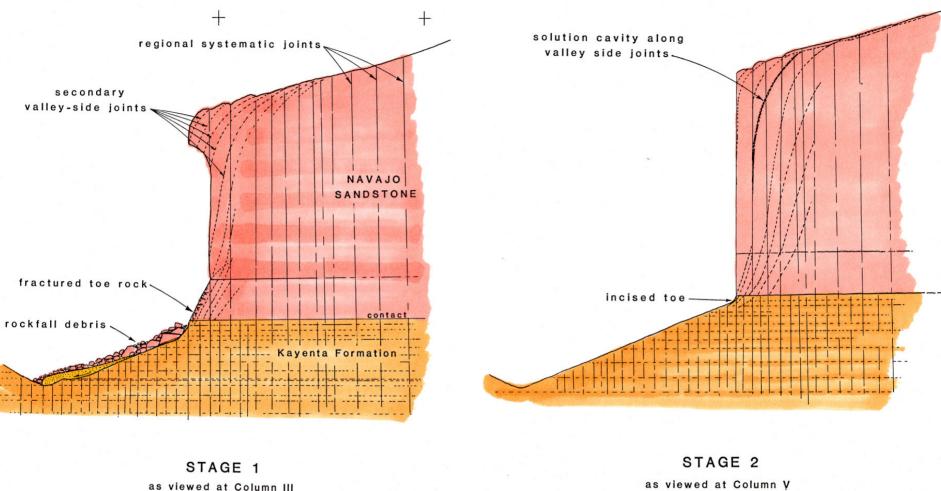
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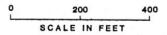


Unravel geomorphic progression – Cliffs tend to retreat in episodic steps. All of these steps are usually discernable at any given site.

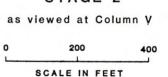


TYPICAL SEQUENCE OF CLIFF RETREAT

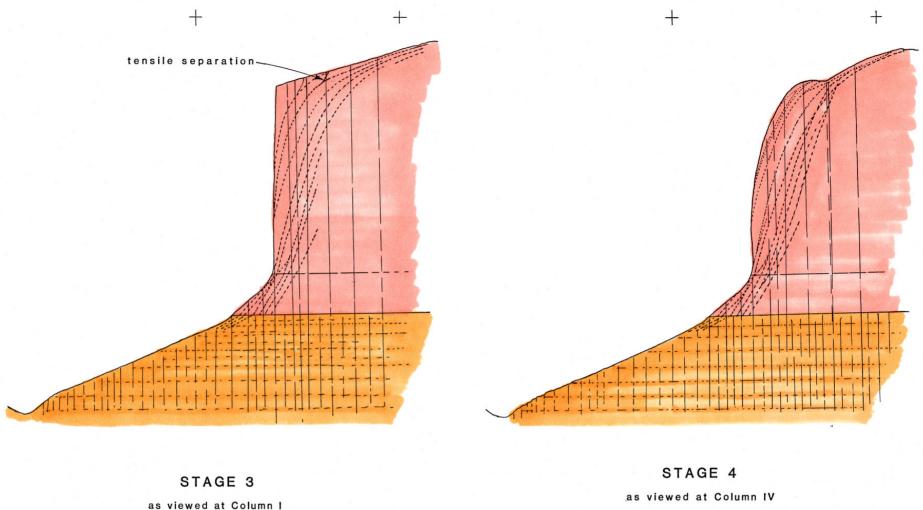




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Stage 1 is gross cliff retreat immediately following a major detachment. After the remaining overhang drops, a completely smooth face forms, arbitrarily designated here as Stage 2.

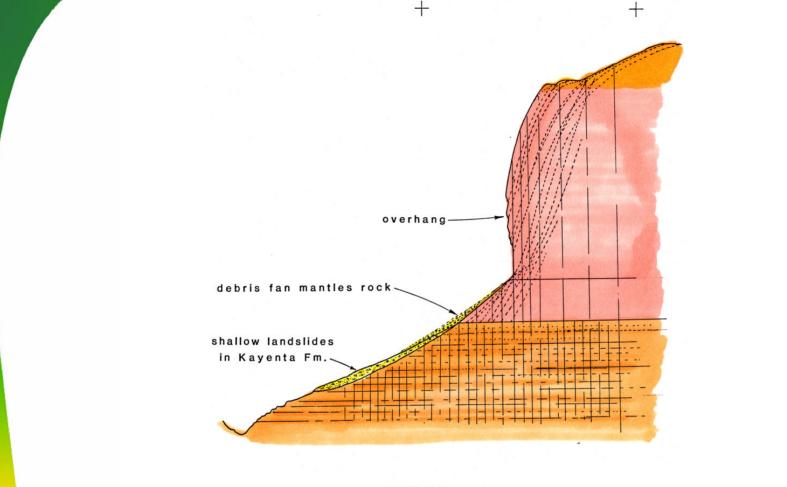


200 400 SCALE IN FEET

200 400 SCALE IN FEET

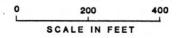
Stage 3 a sloping cliff toe forms, controlled by valley-side joints. In Stage 4 crown blocks have detached along inclined valley-side joints







as viewed at Column III before failure



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Stage 5 is overhanging situation preceding massive detachment of block(s) along inclined valley-side joints

CONCLUSIONS

- Engineering geologists need to unravel the physical agents responsible for shaping the landscape around us
- This requires through background research, patient site mapping, a focused program of subsurface exploration, and a critical analysis of landforms
- We will not identify those geologic features we are not specifically looking for, or with which we have little previous exposure or experience



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